

Melting temperature of diamond at ultrahigh pressure

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Since Ross proposed that there might be ‘diamonds in the sky’ in 1981 (ref. 1), the idea of significant quantities of pure carbon existing in giant planets such as Uranus and Neptune has gained both experimental² and theoretical³ support. It is now accepted that the high-pressure, high-temperature behaviour of carbon is essential to predicting the evolution and structure of such planets⁴. Still, one of the most defining of thermal properties for diamond, the melting temperature, has never been directly measured. This is perhaps understandable, given that diamond is thermodynamically unstable, converting to graphite before melting at ambient pressure, and tightly bonded, being the strongest bulk material known^{5,6}. Shock-compression experiments on diamond reported here reveal the melting temperature of carbon at pressures of 0.6–1.1 TPa (6–11 Mbar), and show that crystalline diamond can be stable deep inside giant planets such as Uranus and Neptune^{1–4,7}. The data indicate that diamond melts to a denser, metallic fluid—with the melting curve showing a negative Clapeyron slope—between 0.60 and 1.05 TPa, in good agreement with predictions of first-principles calculations⁸. Temperature data at still higher pressures suggest diamond melts to a complex fluid state, which dissociates at shock pressures between 1.1 and 2.5 TPa (11–25 Mbar) as the temperatures increase above 50,000 K.

As a result of the importance of high-pressure carbon in both planetary science and inertial confinement fusion (for which high-density carbon is a candidate ablator material for ignition target designs), there are many theoretical calculations of the high-pressure melting curve of diamond, with some predicting a maximum in temperature at 500 GPa (5 Mbar; refs 8–11). Direct temperature measurements are challenging for both static and dynamic high-pressure experiments, so confirmation of melt-curve predictions have been only by inference, never by direct measurement. Under static conditions, equation-of-state data relevant to diamond melting do not extend to pressures much above 50 GPa, and are interpreted as indicating a positive Clapeyron slope, $(\partial T/\partial P)_{\text{melt}} > 0$, for diamond up to 60 GPa (refs 5, 6, 12, 13). Dynamic shock experiments on graphite, which is thought to convert to diamond under dynamic loading, suggest a positive melting slope to at least 140 GPa (refs 14–16). Finally, recent shock experiments on diamond show an increase in density near 600 GPa, suggesting that shocked diamond might potentially melt with a negative melt curve near 600 GPa (refs 17–20).

Here we report the first temperature measurements for shock-compressed diamond at conditions of 0.6–4 TPa (6–40 Mbar) and 8,000–100,000 K. These data reveal the melting curve for carbon up to 1.1 TPa and a complex fluid state between 1.2 and 2.5 TPa.

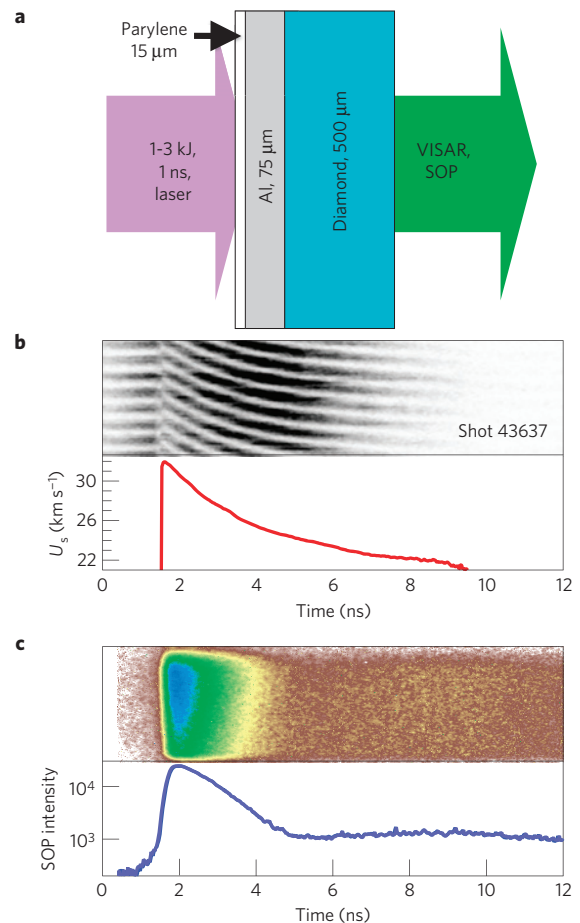


Figure 1 | Experimental configuration and data collected to determine the melting temperature of diamond.

a, Diamond melt target. **b**, Velocity interferometry (VISAR) data for polycrystalline diamond, showing raw fringe data and velocity (U_s) lineout. The pressure falls continuously and smoothly with time as U_s decreases. **c**, SOP data, showing raw data and intensity lineout. The SOP intensity plateaus and rises even as U_s decreases smoothly.

Shock experiments traditionally study a steady-shock Hugoniot characterized by the shock velocity U_s , particle velocity U_p , pressure P , specific volume V , ($=1/\text{density}$) and internal energy E of the shocked material. These five variables are related by the three Rankine–Hugoniot relations, so that the state of the sample under

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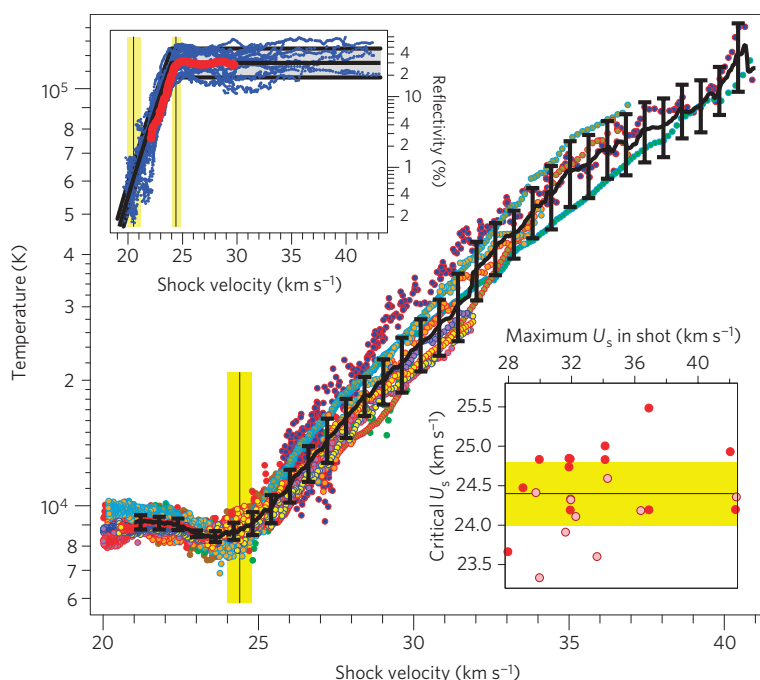


Figure 2 | Temperature versus shock velocity for all nine experiments carried out in the present study. As the shock-front pressure is decaying with time, the highest T, P points correspond to early-time conditions. The upper left inset shows reflectivity measurements (blue) together with data from Bradley *et al.*⁷ (red). The black line is the averaged value. The lower right inset shows the fitted determination of $T-U_s$ kinks (red) and $R_{\lambda=532}$ plateaux (pink) used to find the upper range of the mixed-phase region. The mean and standard deviation of these independent determinations are shown by yellow shaded lines throughout the figure, as is the shock velocity at $R_{\lambda=532} = 0.5\%$ in the upper left inset. The uncertainties shown in the main figure represent the weighted standard deviation of all measurements at each shock velocity as discussed in the Supplementary Information.

shock loading is determined by a single measurement if the P – V – E Hugoniot has been previously determined. Our experiments use an unsupported shock for which the pressure decays with time, while at the shock front P, V, E, U_s and U_p still obey mass, momentum and energy conservation, and hence the Rankine–Hugoniot relations²¹. Thus, the unsupported shock samples a continuum of Hugoniot states as the shock evolves through time and space. We measure the shock velocity U_s in the present experiments, and use the Sesame Hugoniot²² that well represents the compilation of new Hugoniot experiments^{17–20}. In addition, we measure the non-Rankine–Hugoniot variables T , and reflectivity, $R_{\lambda=532}$, of our sample at visible wavelengths ($\lambda = 532$ nm).

We used the OMEGA laser at the University of Rochester to produce shocks in diamond targets by focusing 1 ns pulses with up to 3 kJ of 351 nm light to a flat spot ~ 600 μm in diameter (Fig. 1a). Both single-crystal, natural Type-1a diamond and polycrystalline, synthetic Type-2a diamond discs were used, the latter being formed by chemical vapour deposition and having an average grain size of about 130 μm with little preferred orientation. Any differences due to sample type were well within our experimental uncertainties. Under ambient conditions, the density (ρ_0) and refractive index (n) of the diamond were determined to be $\rho_0 = 3.51(\pm 0.02)$ g cm^{-3} and $n = 2.42(\pm 0.02)$ at $\lambda = 532$ nm.

For most of its transit through the sample, the decaying shock front is reflecting and the instantaneous U_s and $R_{\lambda=532}$ are measured as a function of time using a velocity interferometer (VISAR) (Fig. 1b). $R_{\lambda=532}$ is obtained by comparing the intensity of the VISAR signal from the shock front in diamond with the pre-shock signal from an aluminium reference. Simultaneously, the absolute spectral radiance at two wavelengths, $I(\lambda)$ with $\lambda = 650$ and 450 nm, is measured as a function of time using a streaked optical pyrometer²³ (SOP) and, in combination with $R_{\lambda=532}$, is used to determine T (Fig. 1c; ref. 24) as discussed in the Supplementary Information.

When the unsupported shock enters the diamond at 1.8 ns, $I(\lambda)$ (and T) rises quickly to a peak value (Fig. 1c). These early-time high- P and U_s data correspond to the highest T data at the far right in Fig. 2. Between ~ 2 and 5 ns, $I(\lambda)$ decreases, as expected, owing to the decay of the shock pressure versus time (Fig. 1c). However, between 6 and 10 ns, $I(\lambda)$ actually increases while U_s and P continue to decay. This increase of $I(\lambda)$ was observed in every experiment at $U_s^{\text{crit}} = 24.6(\pm 0.4)$ km s^{-1} (Fig. 2). Similarly, $R_{\lambda=532}$ has a saturated value of $30(\pm 10)\%$ at the initially high shock P and U_s , and then begins to decrease with decreasing U_s at $U_s^{\text{crit}} = 24.3(\pm 0.3)$ km s^{-1} (Fig. 2, upper inset) until the shock front becomes opaque at $U_s \leq 20.5$ km s^{-1} . Combining the two results, we find a significant change in the diamond Hugoniot at a critical shock velocity of $24.4(\pm 0.4)$ km s^{-1} (Fig. 2, lower inset).

If there were no energy sink, the Hugoniot T would decrease systematically as a function of P . Instead, we observe a plateau in shock T between 0.6 and 1.1 TPa (Fig. 3), which we interpret as owing to the latent heat of a phase transition. As only one plateau is observed up to very high T , we identify it with the transition from liquid to solid (corresponding to a two-phase region; ref. 25). Thus, we infer that diamond melts at about 9,000 K and between 0.60 and 1.05 TPa along the Hugoniot. These conclusions are consistent with first-principles molecular-dynamics calculations of the melt curve (Fig. 3; refs 8, 11), the interpretation of the loss of free-surface reflectivity at 550 GPa (ref. 7) and with recent measurements of small U_s – U_p slope changes that also agree with the molecular-dynamics calculations^{8,19}. Our experiments do not have the sensitivity to distinguish the proposed occurrence of the BC8 phase or diamond–BC8–liquid carbon triple point^{8,9,21}.

Reflectivity data for shock-compressed diamond suggest that the fluid phase of carbon is conducting at P and T above 1 TPa and 10,000 K (Fig. 2), in good agreement with first-principles calculations, which predict metallization of carbon in the fluid phase at comparable conditions^{8,26}. $R_{\lambda=532}$ increases with P , from

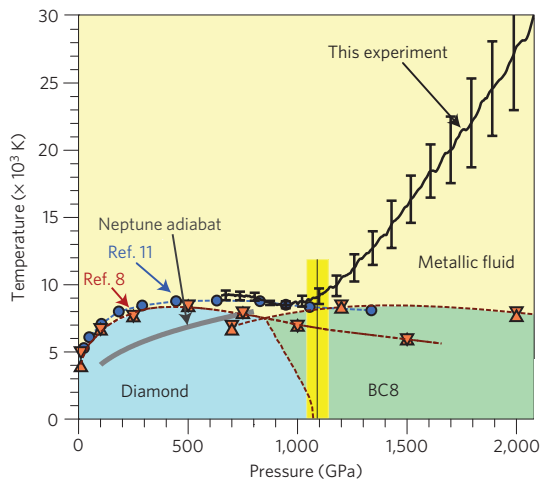


Figure 3 | Temperature versus pressure data compared with simulations.

The black line denotes our new experimental shock temperature data. The blue circles and blue dashed line represent the melt curve from density functional theory calculations from Wang *et al.* that include diamond and liquid phases (ref. 11). The brown triangles that bracket the phase transitions and the brown dashed lines represent similar calculations by Correa *et al.* that include the diamond, BC8 and fluid phases (ref. 8). Our measurements are unable to identify multiple solid phases—be it diamond, BC8 or both—as predicted by the theory. The graphite phase is omitted for clarity. The uncertainties represent the weighted standard deviation of all measurements as discussed in the Supplementary Information.

below our detection threshold of 0.5% at $P < 0.6$ TPa to 30% for $P > 1.1$ TPa. Applying a Drude model, these data give optical depths of ~ 400 nm near 0.6 TPa, falling to ~ 10 nm above 1.1 TPa, as discussed in the Supplementary Information. As atoms undergo more than 15 oscillations and the pressure decay is small during shock transit through an optical depth, the Hugoniot state is expected to be in local thermodynamic equilibrium. Although the transition from mixed-phase to fully solid carbon was not directly determined in the present experiments, both the reflectivity saturation in the pure fluid phase and the theoretical finding that the fluid is metallic⁸ suggest that the onset of melting be identified with the onset of reflectivity. The average U_s at $R_\lambda = 532 = 0.5\%$ (our detection limit) is $U_{sR=0.5\%} = 20.5 \pm 0.6$ km s⁻¹ (0.60 ± 0.06 TPa and $9,200 \pm 800$ K).

Interpreting the diamond Hugoniot as following the melting curve from 0.6 to 1.05 TPa allows us to estimate the Clapeyron slope, $dT_m/dP = \Delta V_m/\Delta S_m = -2.6(\pm 0.6)$ K GPa⁻¹ for the melting curve over this range. The entropy change, ΔS_m , is positive on melting, which implies that the volume change on melting is negative. A negative Clapeyron slope over this pressure range is in good agreement with the predictions of theory (Fig. 3), and a decrease in volume as the Hugoniot crosses the melting temperature of diamond is also supported by recent shock-compression experiments^{17–20}.

Using our Hugoniot temperature measurements, we can extract the specific heat²⁷, as discussed in the Supplementary Information. Just before melting, the sample is at a temperature much higher than the Debye temperature of diamond ($\sim 2,200$ K), so the specific heat should be close to the Dulong–Petit limit of $3Nk_B$ (k_B is Boltzmann’s constant). In the fluid, the specific heat shows a broad peak with $C_V \approx 4.6k_B$ between 10,000 and 30,000 K (Fig. 4), implying an excess energy sink. At higher temperature, C_V eventually approaches $3Nk_B$ at about 70,000 K after first dropping below the Dulong–Petit limit. We consider it unlikely that this peak in C_V is due to a change in ionization, because the reflectivity remains constant at about 30% over this entire P – T range (Fig. 2, upper inset). High values of specific heat are similarly observed just above the melt temperatures

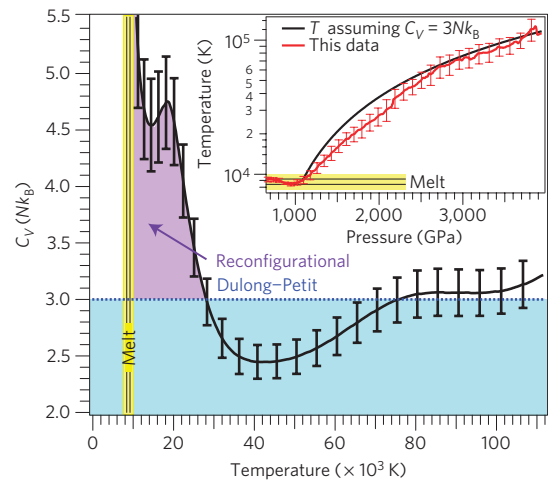


Figure 4 | Specific heat versus temperature showing the Dulong–Petit limit, $C_V = 3Nk_B$, which should hold for the solid below the melting temperature, and the peak attributed to atomic reconfiguration above the melting temperature.

The uncertainties represent the standard deviation of results of a variety of assumptions including variations in the shock Hugoniot and the Gruneisen parameter discussed in the Supplementary Information. The inset compares temperature versus pressure from our measurements (red line) with the values calculated assuming a constant specific heat (black line). The temperatures for a hypothetical fluid obeying the Dulong–Petit limit are well above our uncertainties for pressures of 1.2–2.4 TPa. The temperature uncertainties are identical to those in Figs 2 and 3.

of quartz, fused silica, many alkali halides, silicon and germanium, and have been interpreted in terms of dissociation of a short-range ordered fluid. The high specific heat of Si and Ge is correlated with an anomalous liquid structure factor, and there is an explicit relationship between the structure factor and specific heat for a liquid described by pairwise interactions (equation 4.3.24; ref. 28). Thus, the observed peak in C_V is probably due to a reconfiguration of atomic packing, from a partially bonded complex fluid to an atomic fluid above 60,000 K. (References are presented in the Supplementary Information.)

Finally, estimates of the carbon mass fraction in Neptune and Uranus range from 11 to 17% in the P – T region where CH₄ has been observed to dissociate^{1–3,29}. It is likely that much of this atomic carbon has gravitationally settled into a carbon-rich phase near the core of these planets^{1,3}. Along with our melting temperatures, the model adiabat shown in Fig. 3 suggests that pure carbon could be solid at all depths, although it is possible that a hotter deep interior may exist, dynamically isolated from the outermost atmosphere by a stably stratified, non-convecting region within Uranus and Neptune^{4,30}. In such warmer conditions, pure carbon would exist in the liquid metallic state, settling out of the mantle to form a fluid outer core that helps sustain the planetary magnetic field. In either case, the high sound velocity of a diamond-rich layer at intermediate depths could influence planetary normal modes that may be observable in the future.

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Author contributions

J.H.E., D.G.H., P.M.C., D.K.B. and T.R.B. designed and carried out the experiments. J.H.E., D.G.H., P.M.C. and R.S.M. carried out the analysis; J.E.M. and T.R.B. calibrated the SOP. R.J. and G.W.C. administered the experiment and J.H.E., R.J. and G.W.C. wrote the manuscript.

Additional information

The authors declare no competing financial interests. Supplementary information accompanies this paper on www.nature.com/naturephysics. Reprints and permissions information is available online at <http://npg.nature.com/reprintsandpermissions>. Correspondence and requests for materials should be addressed to J.H.E.